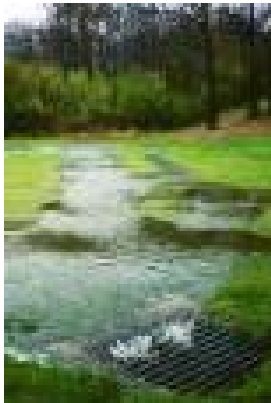


# Lecture 16, Water, Humidity, Pressure and Trace Gases, Part 1

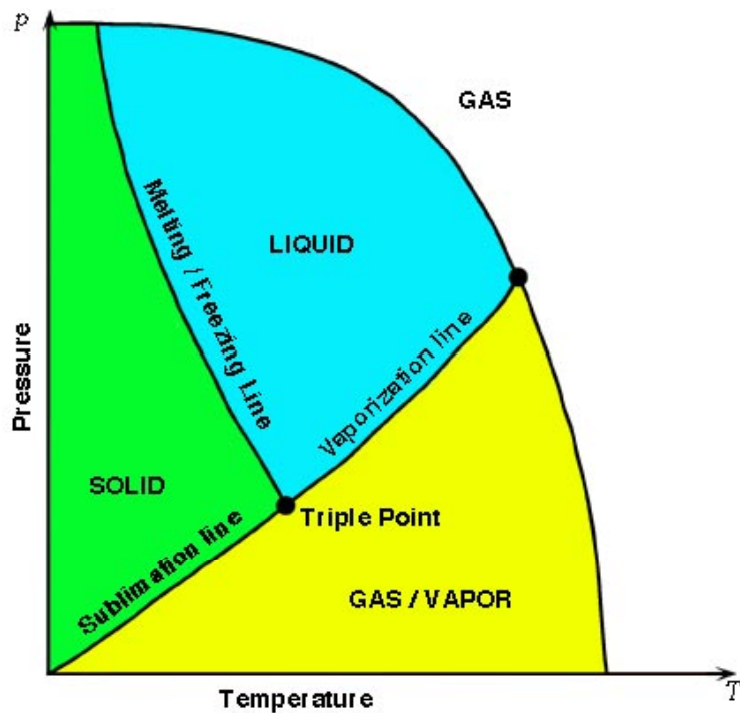
- Physical and chemical properties of water
- Chemical Potential of Water
- Gas Laws
  - absolute humidity
  - relative humidity
  - virtual temperature
  - Saturation Vapor Pressure
  - Clausius-Clapyeron Equation
- Rainfall/Drought

# Water

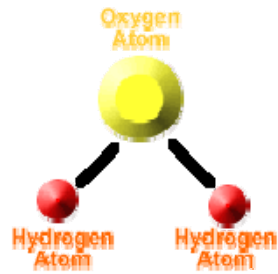
- Gas
- Liquid
- Solid



## Triple Point for Water



How can we Exploit the Triple Point Curve to prevent water from condensing on our air sampling tube?



# Properties of Water

Property	Value
molecular weight	18 g mole <sup>-1</sup>
melting point	273.15 K
boiling point	373.15 K
latent heat of vaporization	2.442 MJ kg <sup>-1</sup> or 44.00 kJ mol <sup>-1</sup> at 20 C
latent heat of sublimation	2.826 MJ kg <sup>-1</sup> or 51.00 kJ mol <sup>-1</sup> at 0 C
dielectric constant	80
thermal conductivity	0.599 W m <sup>-1</sup> at 20 C
heat capacity of water	4182 J kg <sup>-1</sup> at 20 C
molecular diffusivity, water in air	2.42 10 <sup>-5</sup> m <sup>2</sup> s <sup>-1</sup> at 20 C
Density	1.000 kg m <sup>-3</sup> at 4 C

## Properties of Water, f(T)

temperature	density	latent heat of evaporation	kinematic viscosity
°C	Mg m <sup>-3</sup>	kJ mol <sup>-1</sup>	mm <sup>2</sup> s <sup>-1</sup>
0	0.99987	45.0	1.79
4	1.0000	44.8	1.57
10	0.99973	44.6	1.31
20	0.99823	44.1	1.01
30	0.99568	43.7	0.80
40	0.99225	42.8	0.66

# Chemical Potential of Water

**Chemical potential** quantifies  
the driving force for movement of water between two locations

The chemical potential of water is related to the amount of change  
in the Gibbs free energy of the system, subjected by pressure,  
temperature and minor constituents  
, e.g., salts, .

It is a function of its chemical concentration, pressure, electrical potential  
and gravity

The chemical potential has units of energy ( $\text{J mol}^{-1}$ ).

## Water Potential

The **potential energy** of water is related to the difference between its chemical potential ( $\mu_w$ ) and a reference state ( $\mu_{w0}$ ):

$$\psi = \frac{\mu_w - \mu_w^o}{V_w}$$

By convention, Water potential is normalized by  $V_w$ , the partial molal volume of water ( $18.05 \cdot 10^{-6} \text{ m}^3 \text{ mol}^{-1}$ ), giving it units of **Pressure**.

The **total water potential** of a system consists of the sum of water potentials associated with

**Turgor (pressure), osmotic, matrix and gravitational forces**

$$\psi = \psi_p + \psi_\pi + \psi_m + \psi_g$$

Units, Pressure: Pa

- **Turgor (pressure) potential**
  - is related to the hydrostatic pressure, as when someone is blow on or sucking on straw that is inserted in a reservoir of water. Its sign can be positive or negative.
- **osmotic potential**
  - The presence of solutes reduces the activity of water.
- **matric potential**
  - interactions between water and solid surfaces act to reduce the activity of water.
- **gravity Potential**
  - gravitational force is a function of the density of water, the acceleration due to gravity and the height of the water reservoir above or below a reference height:

$$\psi_g = \rho_w gh$$

Water Potential of atmospheric humidity is function of its mole fraction

$$\psi_p = \frac{R \cdot T}{V_w} \ln\left(\frac{e_a}{P}\right)$$

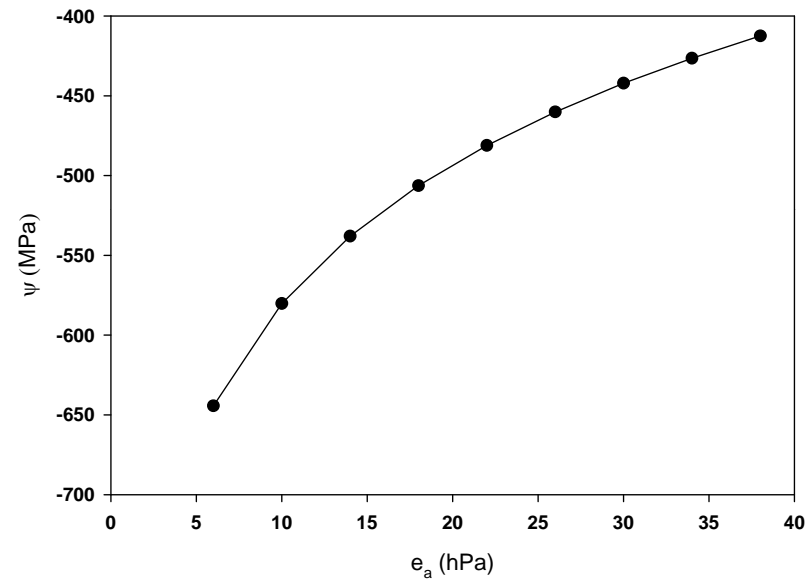
R: Universal gas constant

T: absolute temperature

$e_a$ : vapor pressure

P: pressure

$V_w$ : molal volume of water,  $18.05 \cdot 10^{-6} \text{ m}^3 \text{ mole}^{-1}$



## Atmospheric Gas Composition

constituent	percent by volume	percent by mass	molecular weight (g mole <sup>-1</sup> )
N <sub>2</sub>	78.091	75.5	28.02
O <sub>2</sub>	20.95	23.1	32.00
Ar	0.930	1.3	39.94
CO <sub>2</sub>	0.036	0.05	44.01

## Charles Law and Boyle's Law

pressure, volume and temperature are inter-related.

$$PV = nRT$$

R is the universal gas constant,  $8.3144 \text{ J mol}^{-1} \text{ K}^{-1}$ ,

T is absolute temperature (K),

P is pressure (Pa,  $\text{N m}^{-2}$ ),

V is volume

n is the number of moles

the molar volume of air  
(volume occupied by one mole of air,  $n=1$ )  
for standard pressure (0.1013 MPa, 101.3 kPa, 1013 mb)  
and temperature (273.15 K),

$$V = \frac{nRT}{P}$$

22.41 liters mole<sup>-1</sup> or 0.02241 m<sup>3</sup> mol<sup>-1</sup>.

## Gas Law used by Meteorologists

$$P\alpha = \frac{RT}{m} = \frac{\rho RT}{m}$$

$m$  is the molecular weight of a gas (g mol<sup>-1</sup>)

$\rho$  is the mass density (g m<sup>-3</sup>)

$\alpha$  is 1/ $\rho$ .

## Partial Pressure Law

$$P = p_{N_2} + p_{O_2} + p_{Ar} + p_{CO_2} + p_{H_2O} \dots$$

$$p = \frac{R}{m} \rho T$$

$$\frac{\rho_a}{m_a} + \frac{\rho_v}{m_v} + \frac{\rho_c}{m_c} + \dots = \frac{p}{RT}$$

# Air Density

$$\frac{\rho_a}{m_a} + \frac{\rho_v}{m_v} = \frac{p}{RT}$$

T	dry	Saturated
°C	Kg m <sup>-3</sup>	kg m <sup>-3</sup>
0	1.292	1.289
5	1.269	1.265
10	1.246	1.240
15	1.225	1.217
20	1.204	1.194
25	1.183	1.169
30	1.164	1.145
40	1.128	1.096



Does a baseball fly less far on a humid day?



**mass gas density** (grams of a gas molecule per unit volume)

$$\rho_c = \frac{n_c m_c}{V} = \frac{p_c m_c}{RT}$$

$n_c$  is the number of moles

$V$  is volume

$R$  is universal gas constant

$m_c$  is the molecular mass ( $\text{g mole}^{-1}$ ) of the compound.

**molar gas density** (moles of a gas molecule per unit volume)

$$c = \frac{n_c}{V} = \frac{p_c}{RT} = \frac{\rho_c}{m_c}$$

$n_c$  is the number of moles

$V$  is volume

$m_c$  is the molecular mass (g mole<sup>-1</sup>) of the compound.

**Mass fraction** (mass per unit mass of air)

$$M_c = \frac{\rho_c}{\rho_a} = \frac{\chi_c m_c}{m_a}$$

$m_a$  is the molecular weight of dry air ( $\text{g mol}^{-1}$ )

$m_c$  is molecular weight of trace gas

$\rho_a$  is the mass density of dry air ( $\text{g m}^{-3}$ )

$\rho_c$  is the mass density of trace gas air ( $\text{g m}^{-3}$ )

**Mole fraction,**

the number of moles of a trace gas divided by the total number of moles present in the mixture

Moist air

$$C_c = \frac{n_c}{n} = \frac{m_a \rho_c}{m_c \rho} = \frac{p_c}{p}$$

Dry air

$$C_c = \frac{n_c}{n_d} = \frac{p_c}{p_d} = \frac{p_c}{p_a - p_{H_2O}}$$

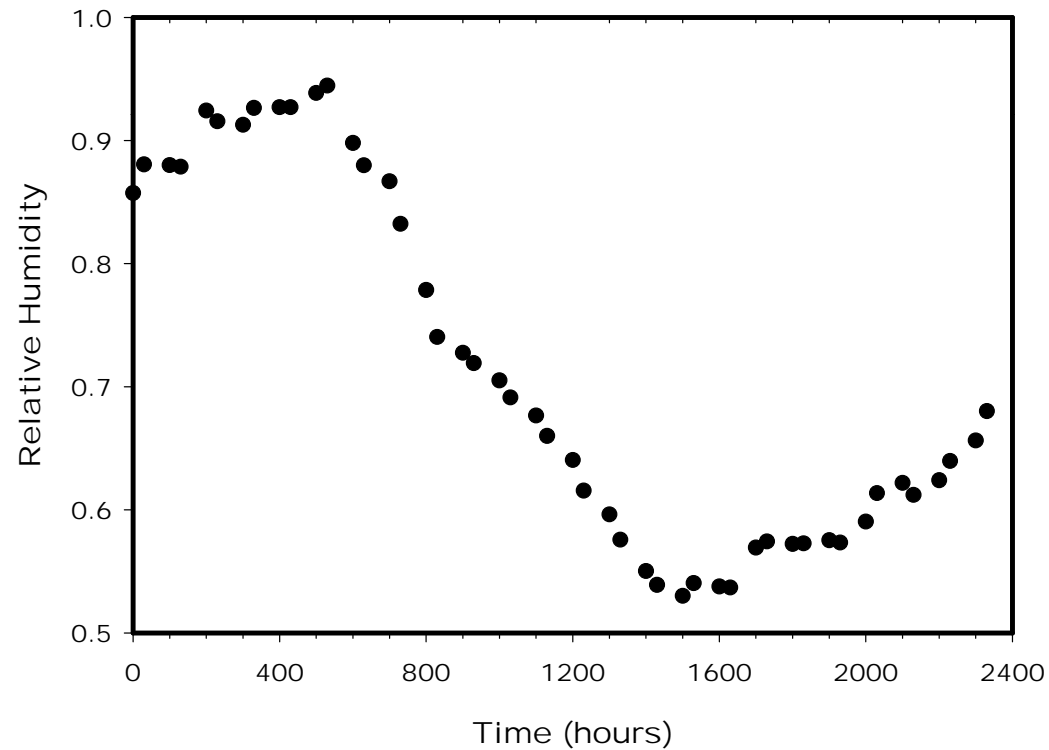
## Relative Humidity, $h_r$ ,

the ratio between the actual ( $e_a$ ) and saturation vapor pressures ( $e_s(T)$ ).

It ranges between zero and 1.0, with one indicating saturation.

$$h_r = \frac{e_a}{e_s(T)}$$

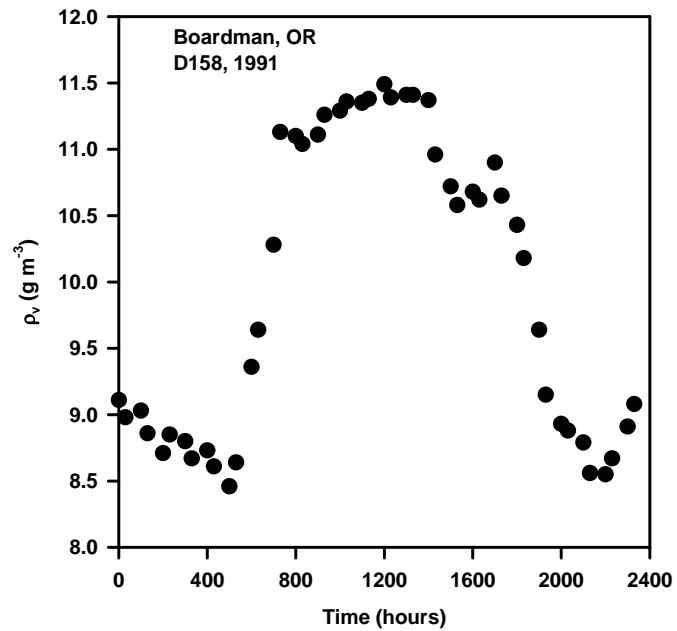
$$e_a = p_{\text{H}_2\text{O}}$$



**Absolute humidity** or vapor density ( $\text{g m}^{-3}$ ):

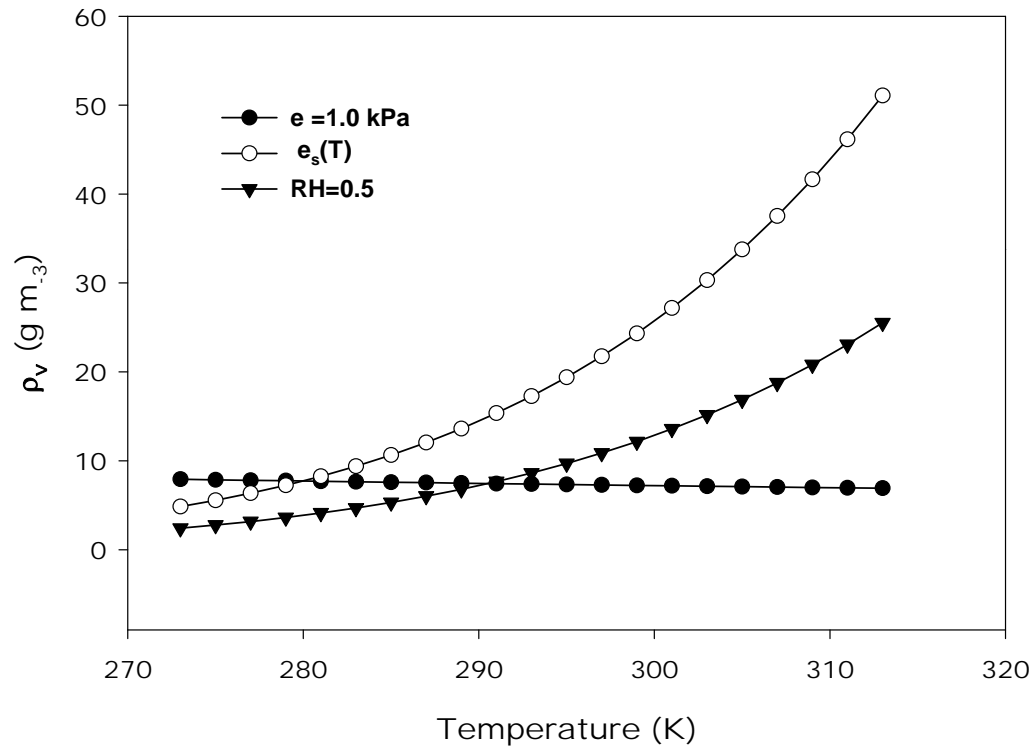
$$\rho_v = \frac{e_a m_v}{RT_k} \quad \rho_v (\text{g} \cdot \text{m}^{-3}) = \frac{2.165 \cdot e_a (\text{kPa})}{T_k}$$

( $R=8.3143 \text{ J mol}^{-1} \text{ K}^{-1}$   $m_v=18 \text{ g mol}^{-1}$ )



Biometeorology, ESPM 129

# Humidity vs Temperature



## Virtual Temperature.

is the temperature **dry air would have**  
**if it had the same density as moist air at the same pressure**

it is related to the **speed of sound**

and

$$v_{sound} = \sqrt{\frac{C_p}{C_v} \frac{P}{\rho}} = \sqrt{\frac{C_p}{C_v} \frac{RT_v}{m_{air}}}$$

**buoyancy of the air.**

$$a = g \frac{(\rho - \rho_{parcel})}{\rho_{parcel}}$$

$$\frac{\rho'}{\rho} = \frac{T_v'}{T_v}$$

Acceleration, a

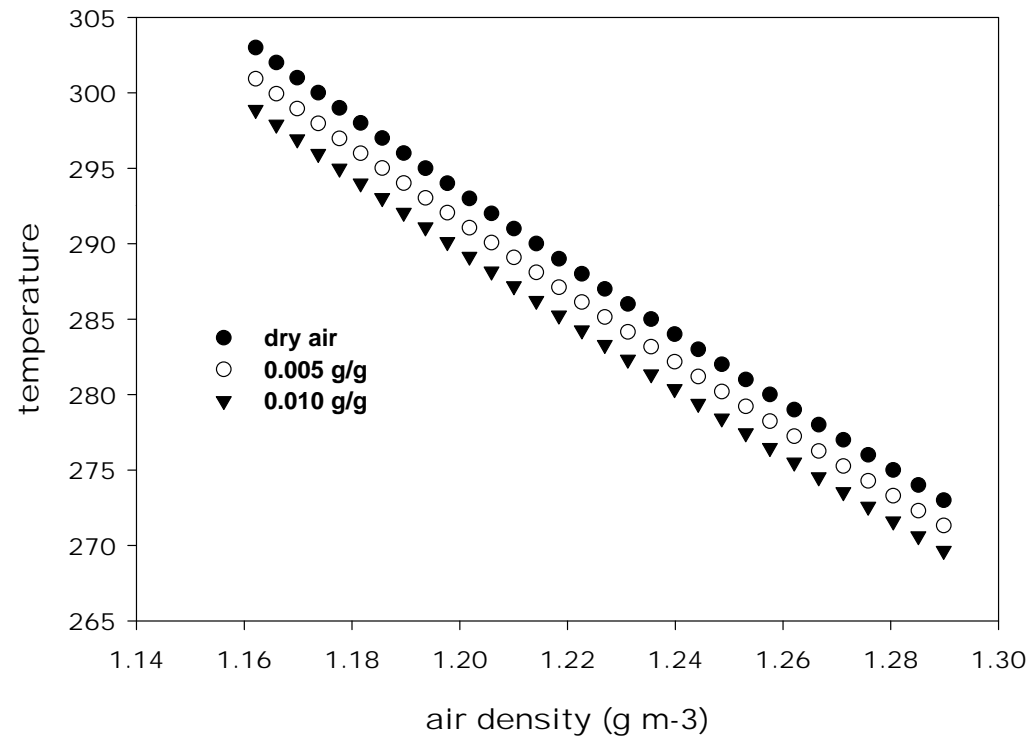
## Derivation: Virtual Temperature

$$P = (P - e)|_{dry} + e|_{moist} = \frac{\rho RT}{m} = RT \left( \frac{\rho_a}{m_a} + \frac{\rho_v}{m_v} \right) = \frac{\rho RT_v}{m_a}$$

$$T_v = T \left( \frac{1 + \frac{m_a \rho_v}{m_v \rho_a}}{1 + \frac{\rho_v}{\rho_a}} \right)$$

$$T_v = T \left( 1 + 0.38 \frac{e}{P} \right)$$

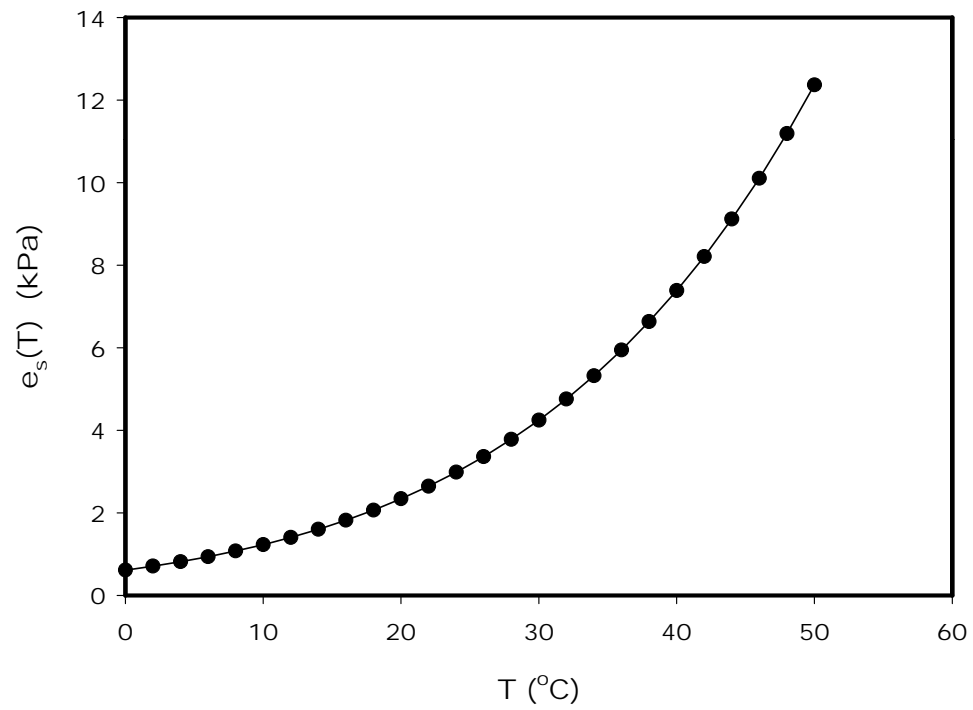
## Air Temperature and Virtual Temperature

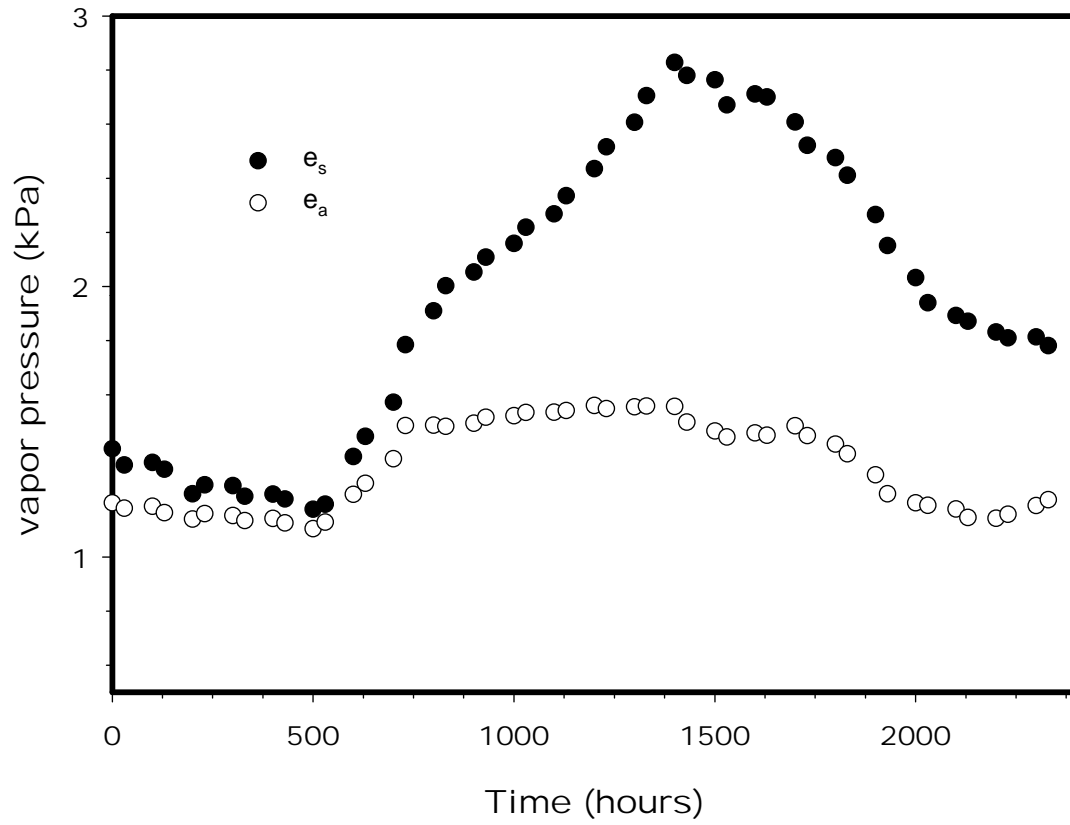


# Saturation Vapor Pressure

- When a pool of water is at constant temperature in a closed container, some water molecules are leaving the liquid and others are condensing and returning to the liquid. Molecules in the head space exert a partial pressure on the system.
- The equilibrium vapor pressure that occurs is called the saturation vapor pressure.
- It is a function of temperature and is independent of pressure.

$$e_s(T_C) = 0.611 \exp\left(\frac{17.502 \cdot T_C}{T_C + 240.97}\right)$$





## Probability of Time between Storms

$$f(\tau) = \lambda \exp(-\lambda\tau)$$

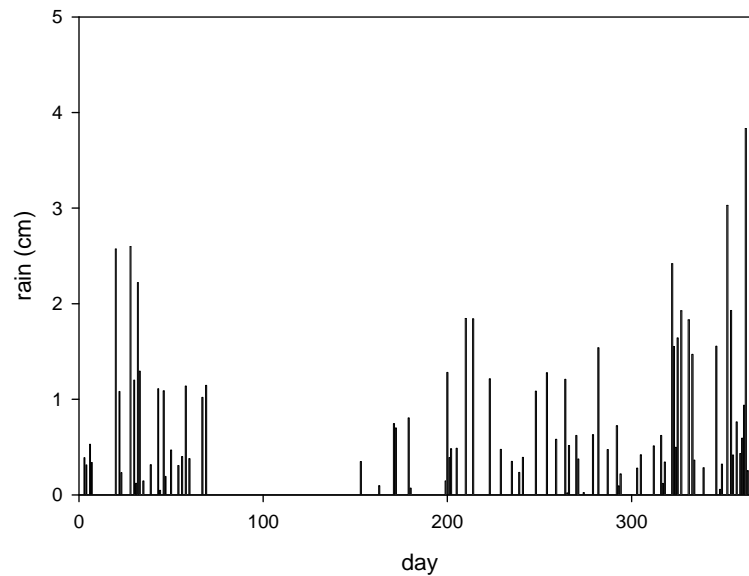
$\lambda$  is the rate of rain (units of rain events per day). It is computed by summing the number of rain days per year and dividing by 365

## Probability of amount of rain per storm

$$f(h) = \frac{1}{\alpha} \exp\left(-\frac{h}{\alpha}\right)$$

$1/\alpha$  is the depth of the rain.  $\alpha$  is computed by summing the amount of rain per year and dividing by the number of rainy days.

Stochastic Rain Fall



# Drought

- Meteorological
  - Periods when precipitation is significantly below the long-term average
- Hydrological
  - When water level in lakes and rivers fall significantly below normal conditions
- Agricultural
  - Occurrence of low levels of plant available water
- Sociological
  - When drought disrupts societies and/or societies disrupt supplies of water, eg via civil war, famine, political decisions

# Drought Metrics

- Palmer Drought Index
- Budyko Aridity Index
- Thornthwaite Index
  - Ratio of between potential and actual evaporation
  - Ratio between precipitation and available energy

# Summary

- The concept of chemical potential energy quantifies the driving force for movement
- of water between two locations, such as in the soil, in a plant and between the soil, plant and atmosphere.
- Matrix, pressure, osmotic, gravitational potential sum to determine the chemical potential of water.
- Saturation vapor pressure is an exponential function of temperature
- The probability of rainfall events and its amount can be computed with a Poisson probability distribution