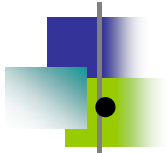


# Leaf Energy Balance, part 2



- IsoThermal Energy Balance
- Critical Limits on Evaporation
  - infinite surface conductance
  - infinite aerodynamic conductance
  - zero surface conductance
  - zero aerodynamic conductance
- Response surfaces, leaf-air temperature differences, Optimal Leaf Form, Size and Shape

# IsoThermal Radiation Balance

$$R_{ni} = Q - \sigma \epsilon T_a^4$$



$$\begin{aligned} R_n &= Q - \sigma \epsilon T_l^4 = \\ Q - (\sigma \epsilon T_a^4 + 4\sigma \epsilon T_a^3 (T_l - T_a)) &= \\ R_{ni} - 4\sigma \epsilon T_a^3 (T_l - T_a) \end{aligned}$$

$$\lambda E = \frac{sR_{iso} + D\rho_a C_p (g_r + g_h)}{(s + \frac{\gamma g_r}{g_w} + \frac{\gamma g_h}{g_w})}$$

$$g_r = \frac{4\epsilon\sigma T_a^3}{\rho C_p}$$



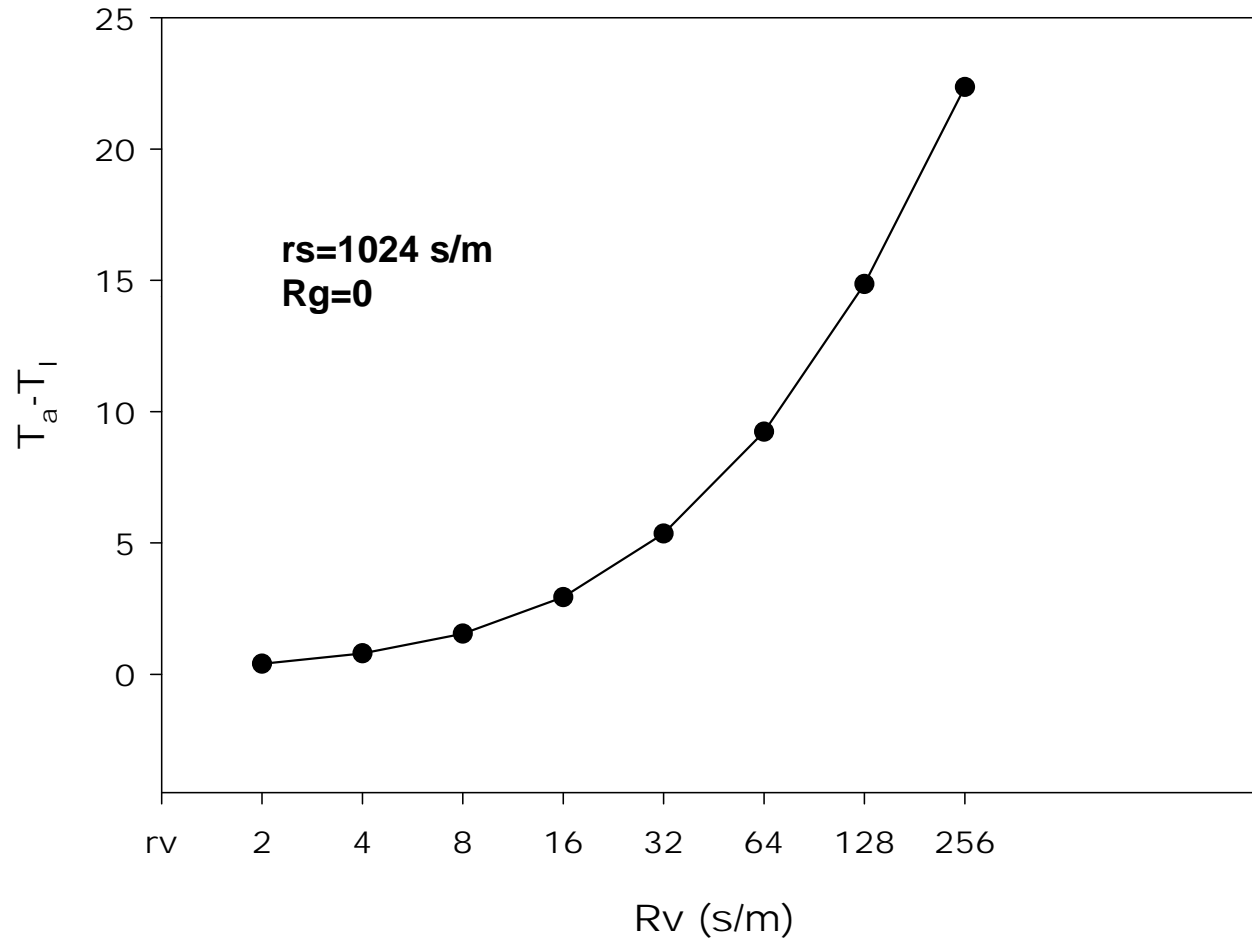
## Nocturnal Evaporation

surface conductance goes to infinity as the surface is wet with condensation

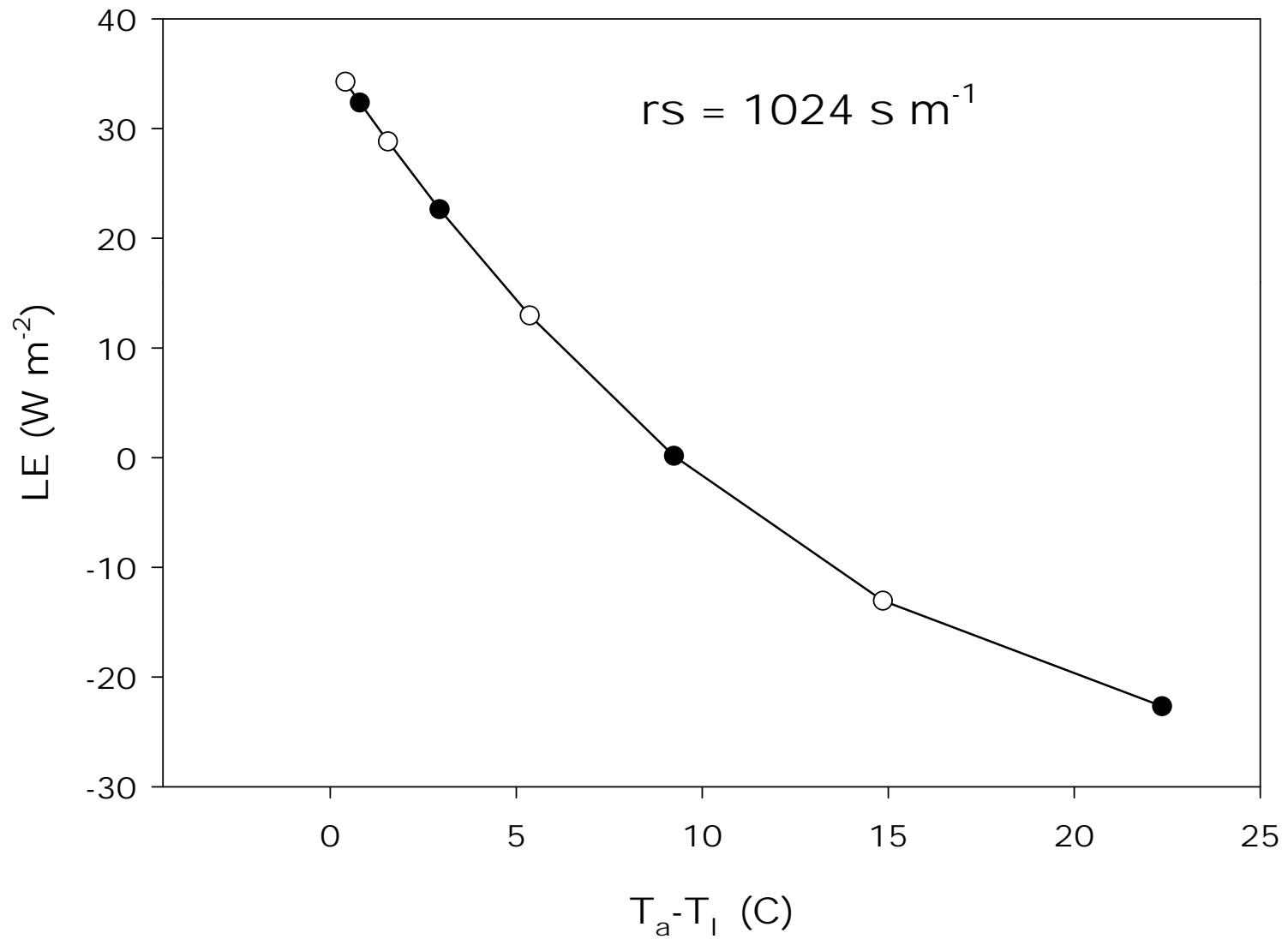
$$\lambda E = \frac{sR_{iso} + D\rho_a C_p (g_r + g_h)}{(s + \gamma(1 + \frac{g_r}{g_{av}}))(1 + \frac{g_{av}}{\infty})}$$

$$\lambda E = \frac{sR_{iso} + D\rho_a C_p (g_r + g_h)}{(s + \gamma(1 + \frac{g_r}{g_{av}}))}$$

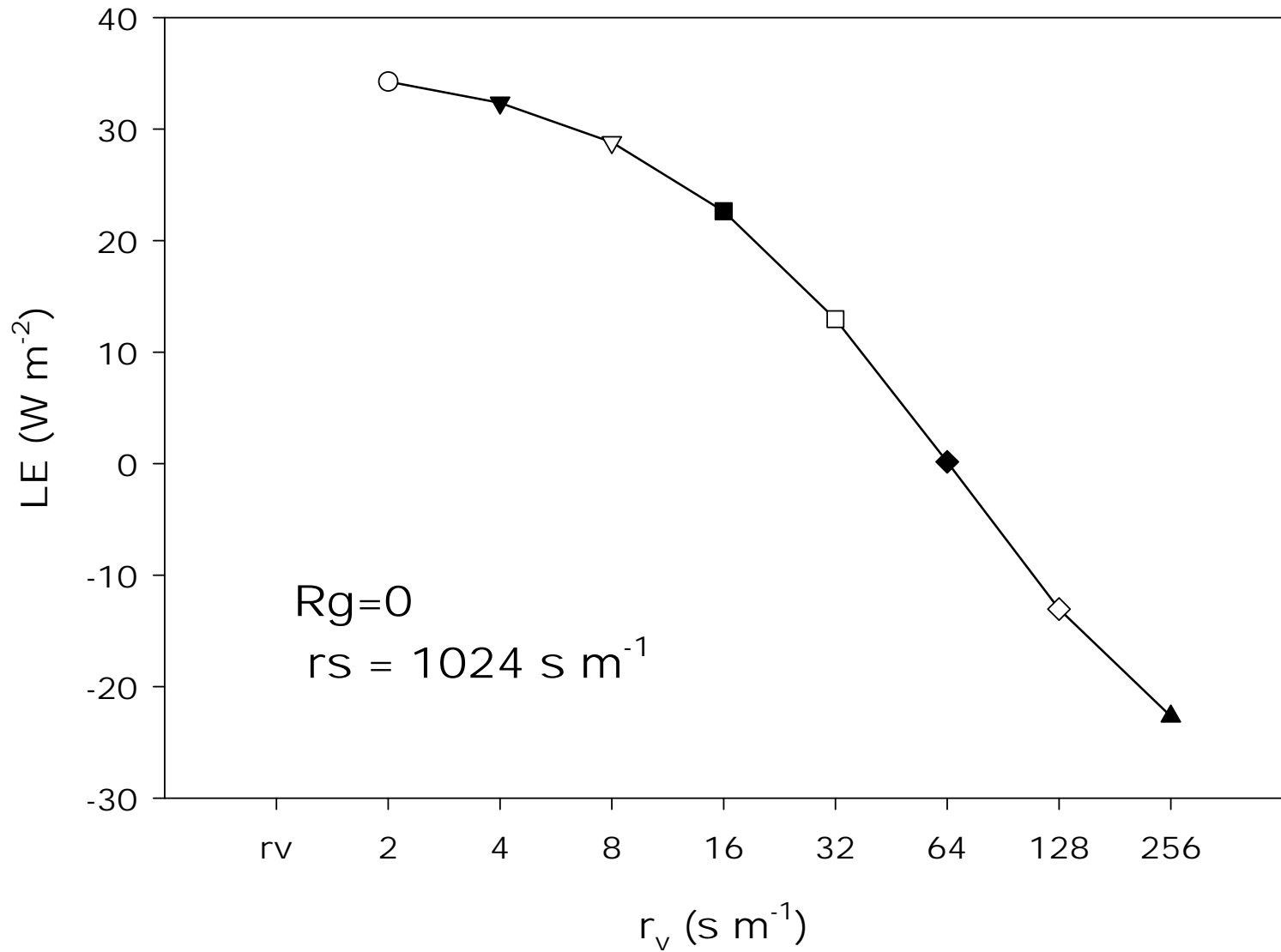
# Leaf-air temperatures at night with closed stomata



# Evaporation/dew formation at night with closed stomata



# Evaporation/dew formation at night with turbulent mixing



## Well-Mixed Case, Infinite Boundary Layer Conductance



$$\lambda E(g_a \rightarrow \infty) = \frac{sR_n + D\rho_a C_p \infty}{(s + \gamma(1 + \frac{\infty}{g_s}))}$$

$$\lambda E_{imposed}(g_a \rightarrow \infty) = \frac{\rho_a C_p}{\gamma} D g_s$$

# Steps



$$\frac{sR_n}{(s + \gamma(1 + \frac{\infty}{g_s}))} = 0$$

$$\lambda E(g_h \rightarrow \infty) \sim \frac{D\rho_a C_p g_h}{(s + \gamma(1 + \frac{g_h}{g_s}))}$$

$$\frac{1}{\{\lambda E\}} \sim \frac{(s + \gamma(1 + \frac{g_h}{g_s}))}{D\rho_a C_p g_h}$$

$$\frac{s}{D\rho_a C_p \infty} + \frac{\gamma}{D\rho_a C_p \infty} + \frac{\gamma g_h}{g_s D\rho_a C_p g_h}$$

$$\frac{1}{\lambda E} \sim \frac{\gamma g_h}{g_s D\rho_a C_p g_h} = \frac{\gamma}{g_s D\rho_a C_p}$$

Still Air Case, boundary layer conductance goes to Zero



$$\lambda E(g_a \rightarrow 0) = \frac{sR_n + D\rho_a C_p 0}{(s + \gamma(1 + \frac{0}{g_s}))}$$

$$\lambda E_{eq}(g_a \rightarrow 0) = \frac{s}{s + \gamma} R_n \quad \text{Equilibrium Evaporation}$$

But, this Equality is physically nonsensical.

How can we expect the surface to exchange water and heat with the atmosphere if there is no transfer mechanism?

Humidity Deficit Must be Infinite in Still Air with Impending Energy



$$D(g_a \rightarrow 0) = \frac{\lambda E (s + \gamma (1 + \frac{0}{g_s}))}{\rho_a C_p 0} - s R_n = \infty$$



Alternatively Consider IsoThermal Radiation Balance

$$\lambda E(g_a \rightarrow 0) = \frac{sR_{iso} + D\rho_a C_p (g_r + 0)}{(s + \gamma(1 + \frac{g_r}{0})(1 + \frac{0}{g_s}))} = \frac{[]}{\infty} = 0$$

It correctly and Physically Reveals Evaporation is Zero in Still Air

## Equilibrium Evaporation

$$\frac{d\lambda E}{dt} = \frac{dE}{dD} \frac{dD}{dt}$$

$$\lambda E_{eq}(t \rightarrow \infty) = \frac{s}{s + \gamma} R_n$$

## Wet Surface



$$\lambda E(g_s \rightarrow \infty) = \frac{sR_n + D\rho_a C_p g_h}{(s + \gamma(1 + \frac{g_h}{\infty}))}$$

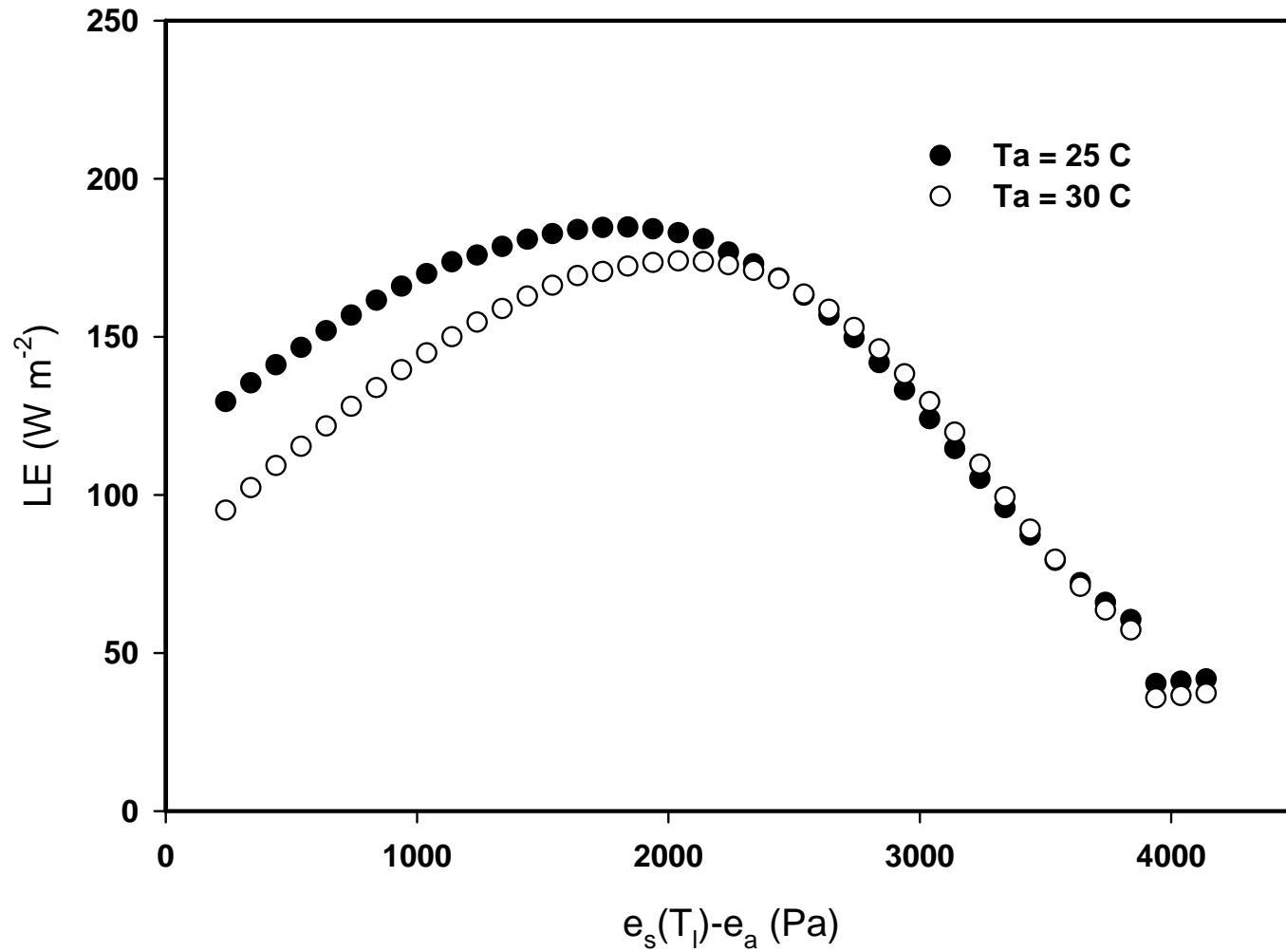
$$\lambda E(g_s \rightarrow \infty) = \frac{sR_n + D\rho_a C_p g_h}{(s + \gamma)}$$

## Coupling Theory

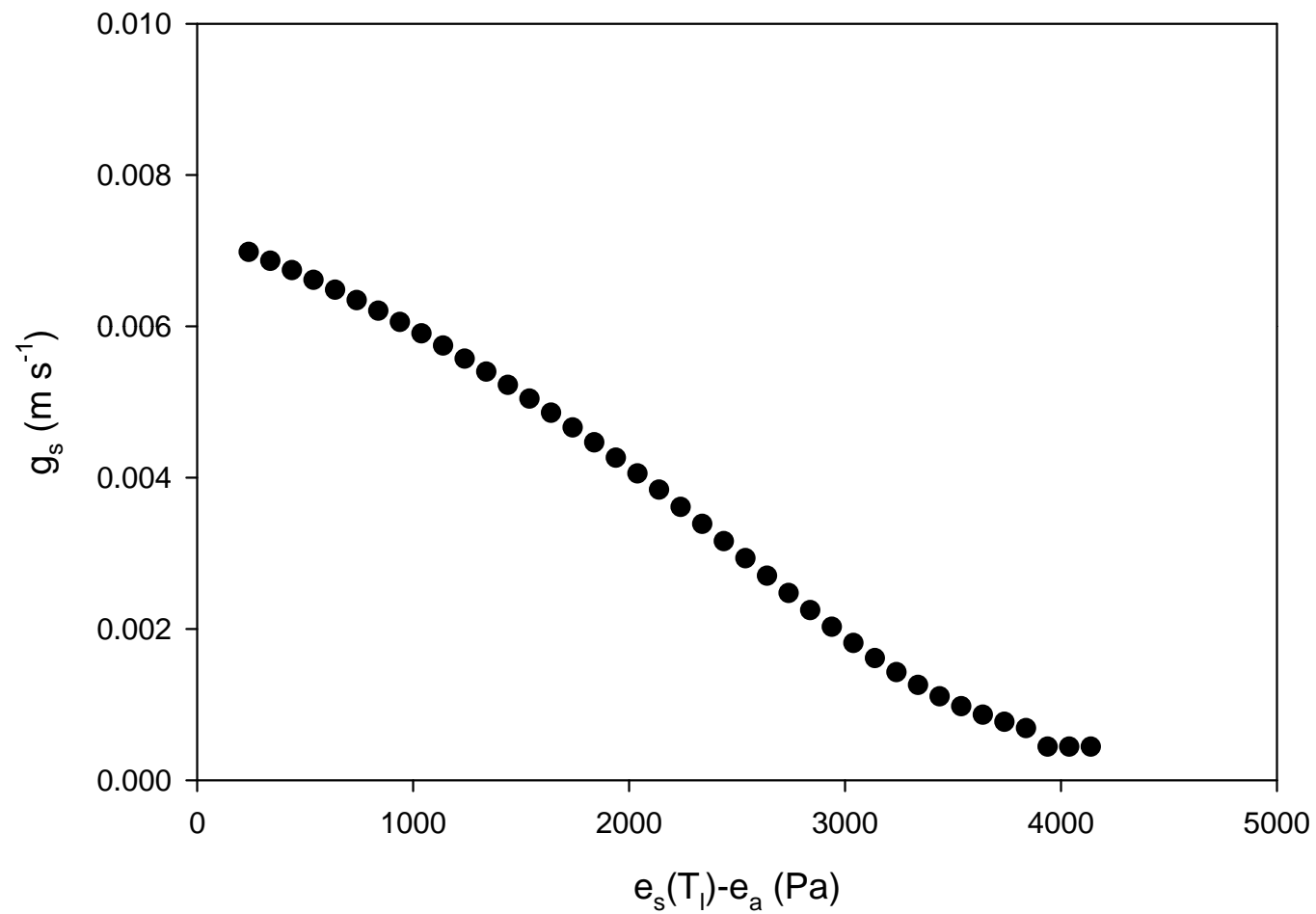
$$dE = (1 - \Omega) \frac{E}{g_s} dg_s \quad \Omega(R_{iso}) = \frac{1 + \varepsilon + \frac{g_r}{g_b}}{1 + \varepsilon + \frac{g_b + g_r}{g_s} + \frac{g_r}{g_b}}$$

species	gs	D (mm)	$\Omega$ (0.2 m/s)	$\Omega$ (5 m/s)
Sitka spruce	0.07	2	0.18	0.03
Beech	0.10	40	0.50	0.10
apple	.21	60	0.50	0.11

# Response of evaporation to changes in vapor pressure differences

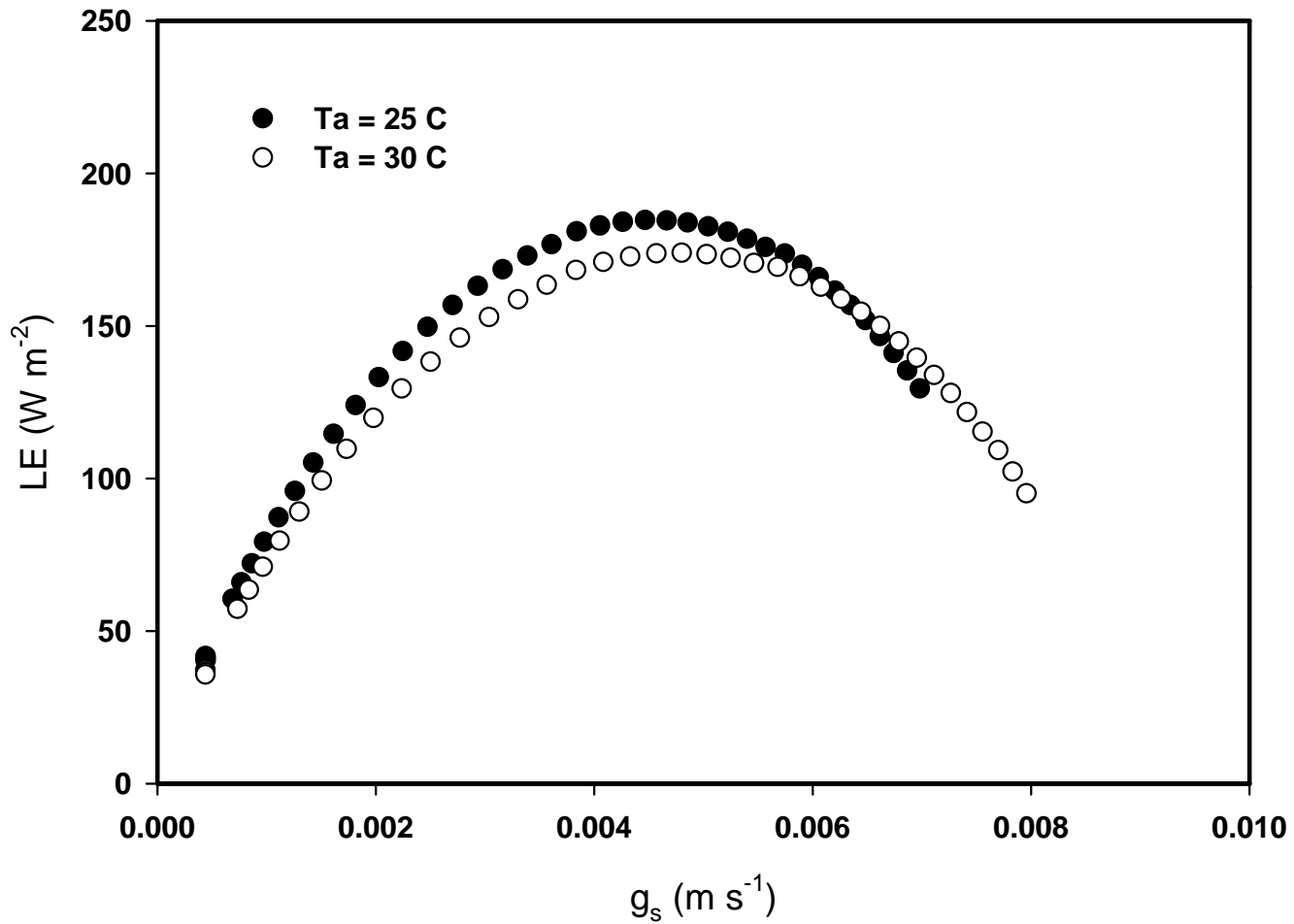


# Response of stomatal conductance to humidity deficits

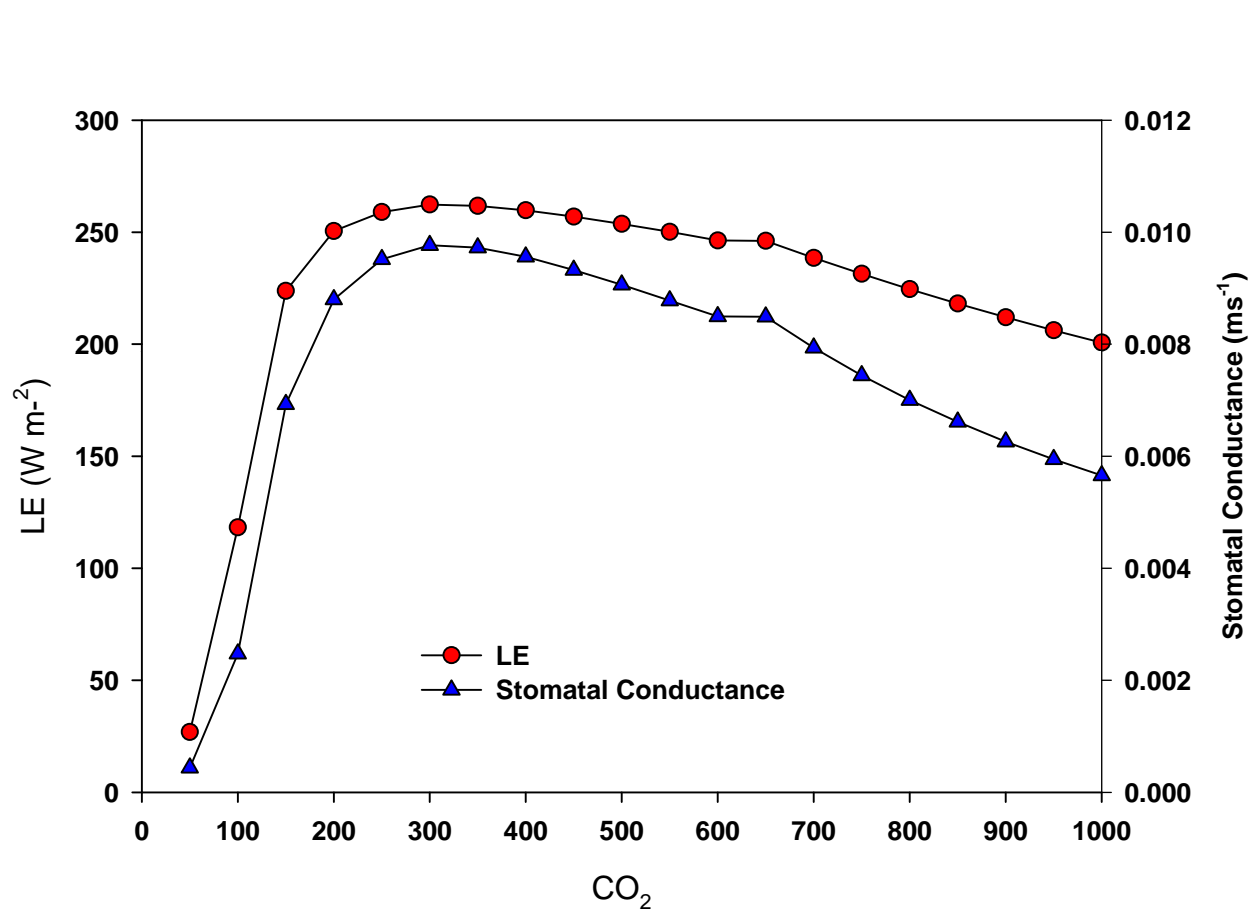


# The relation between LE and stomatal conductances

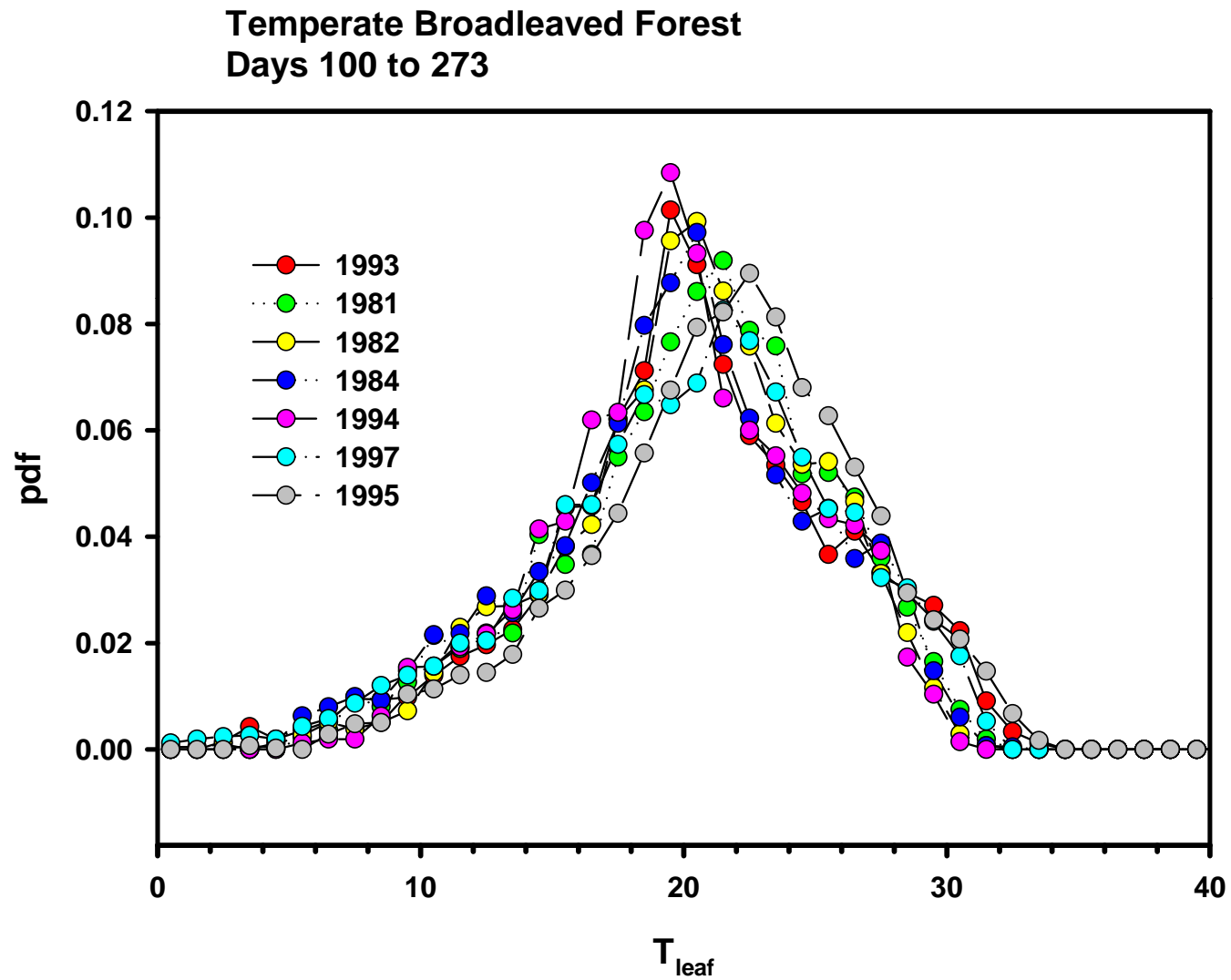
coupled photosynthesis-energy balance-stomatal conductance model  
Increasing humidity;  $R_g = 1000 \text{ W m}^{-2}$ ;  $u = 3 \text{ m s}^{-1}$



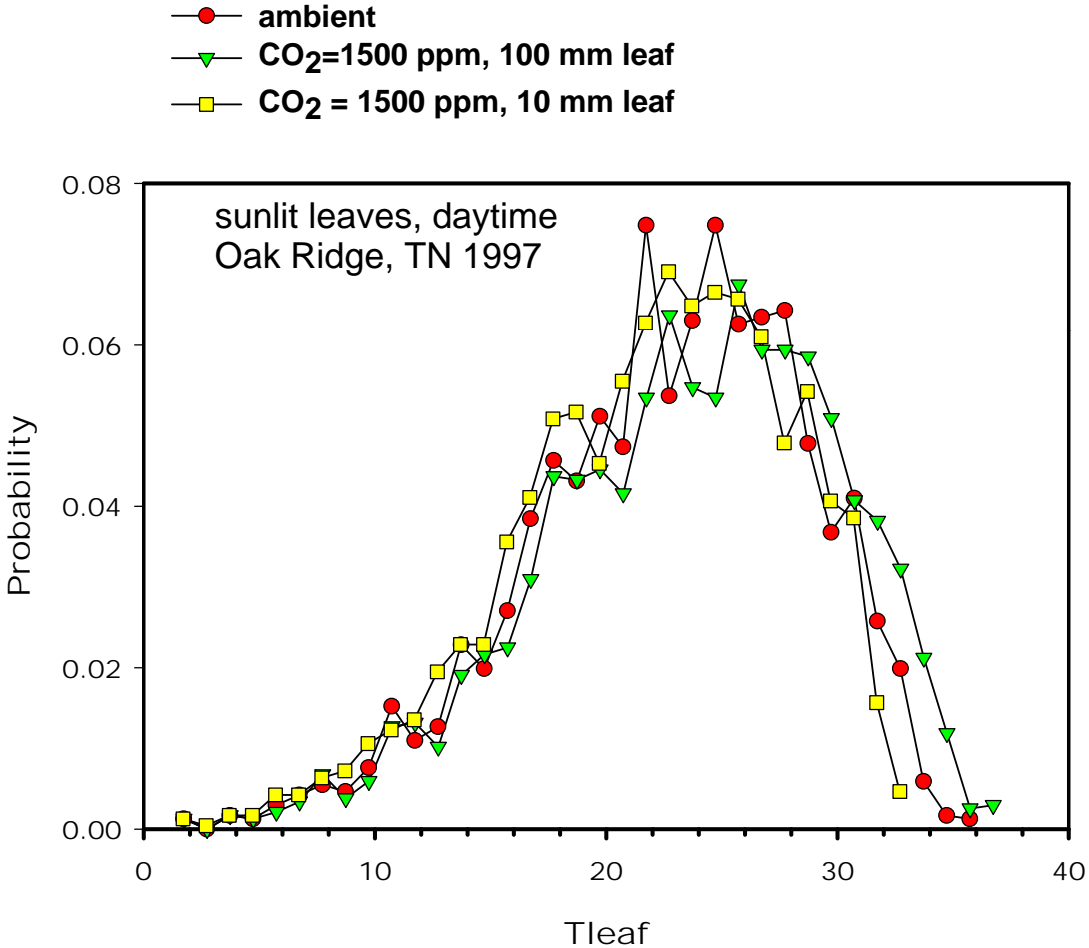
# Transpiration and CO<sub>2</sub>



# Probability density function of mean leaf temperature of a broadleaved forest in Tennessee

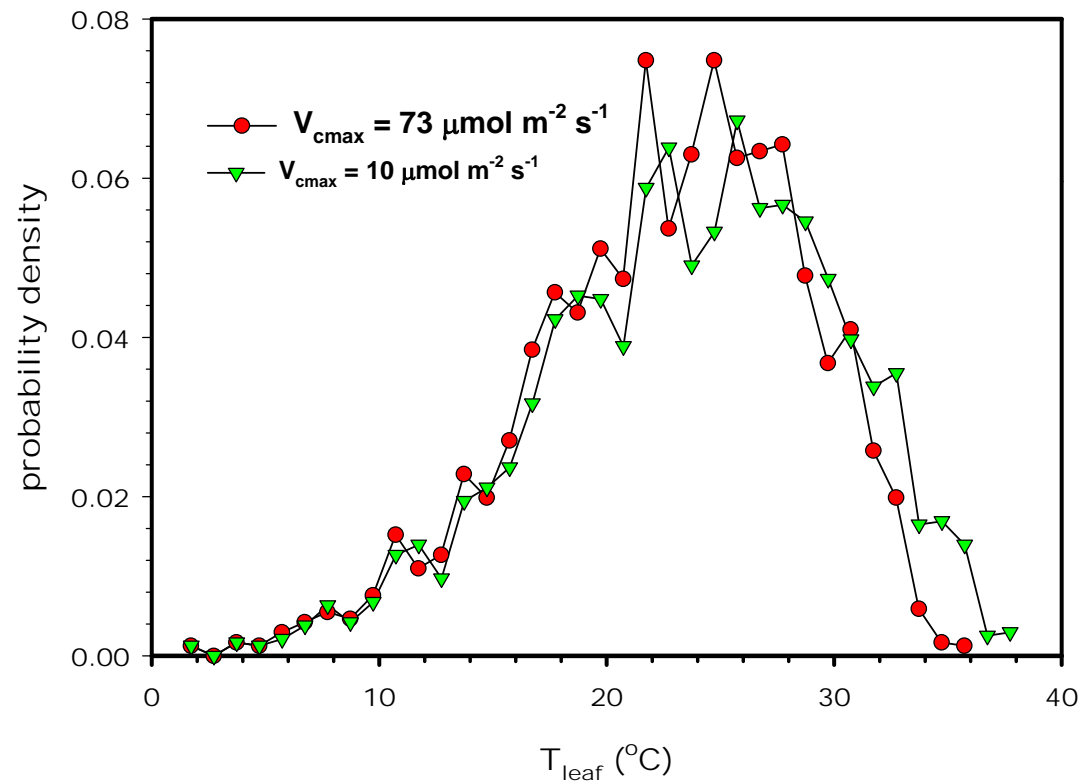


# Theoretical calculations of leaf size based on climate scenarios

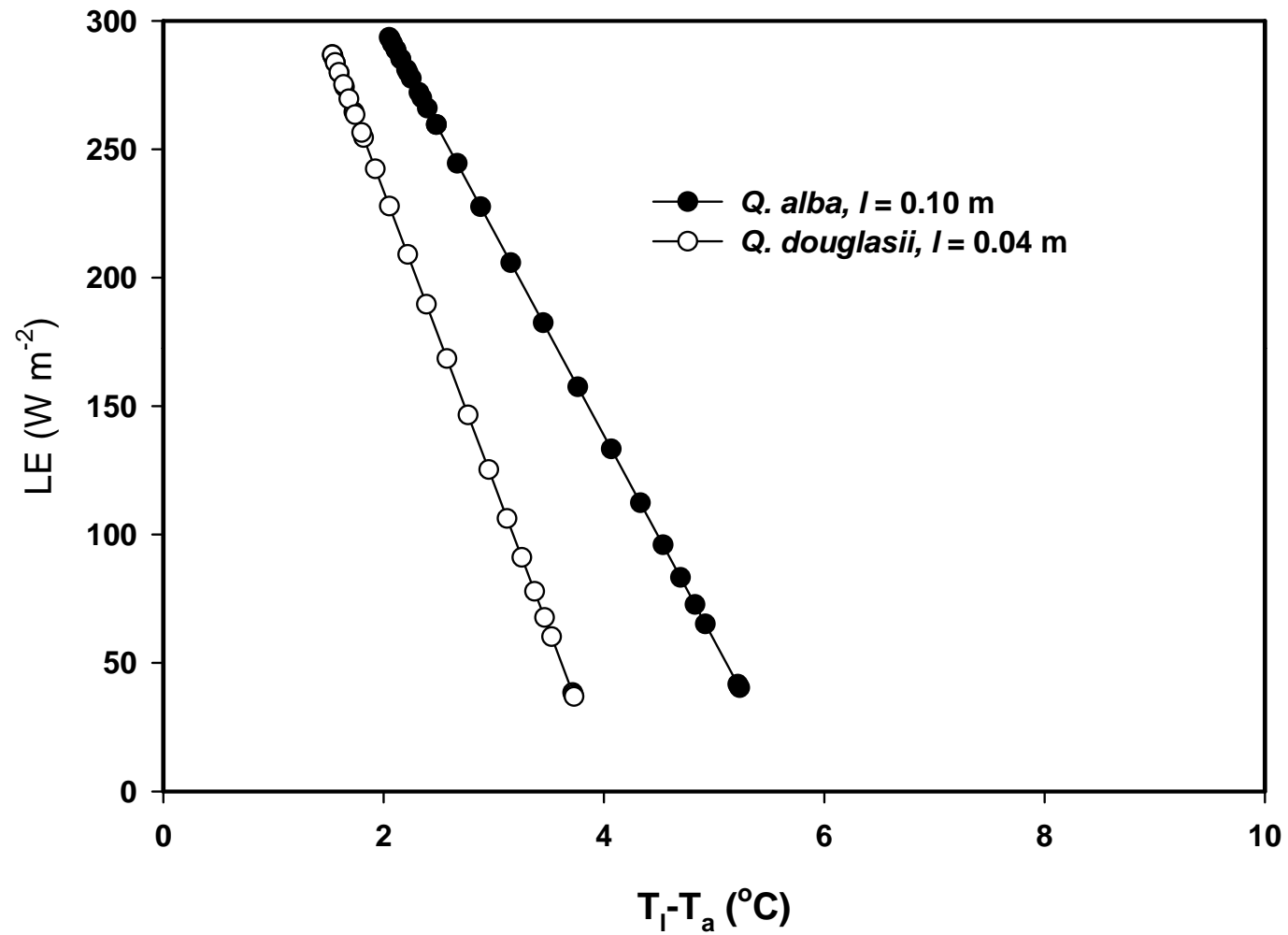


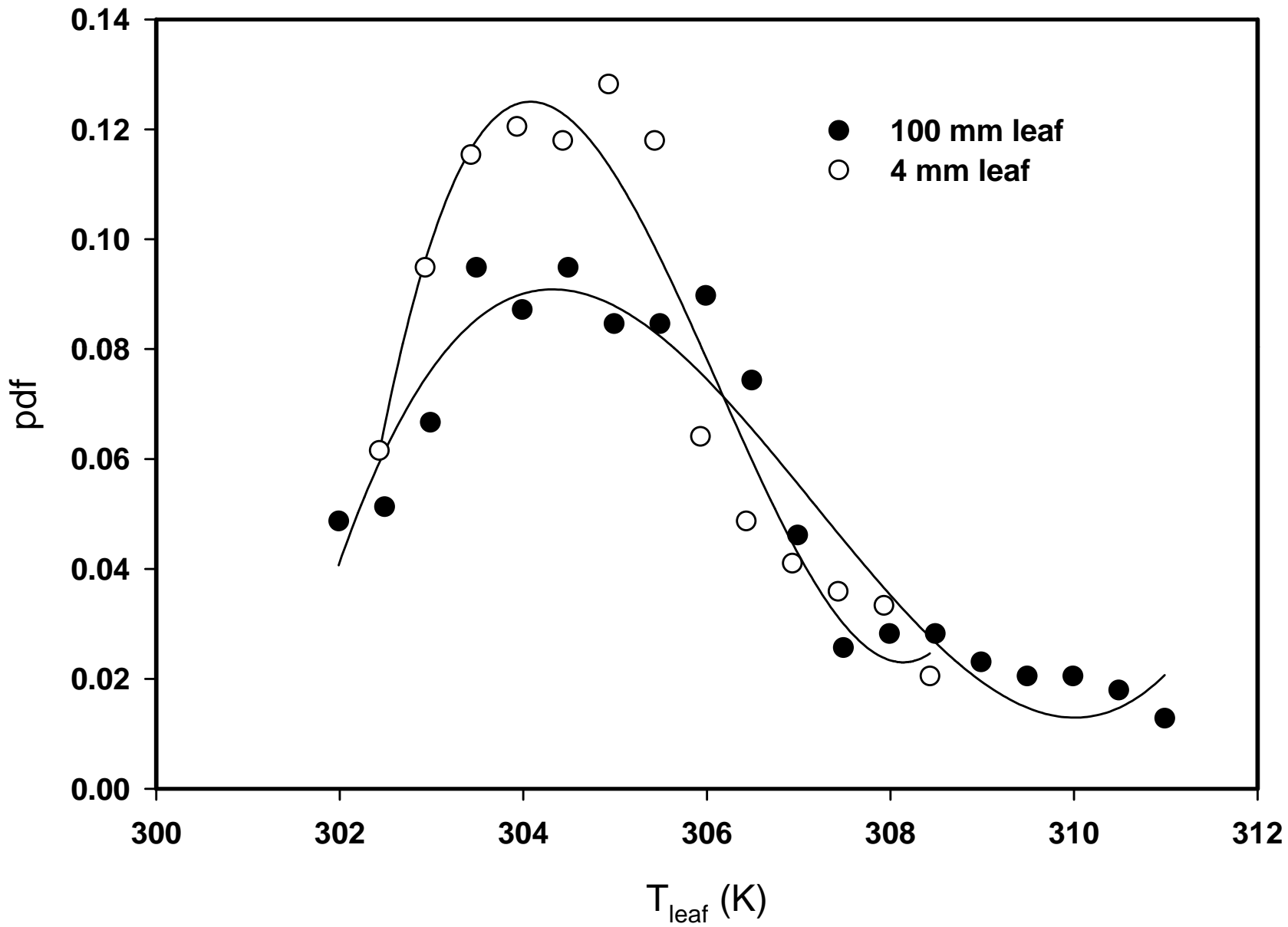
# Role of leaf photosynthetic capacity and leaf temperature

Temperate Deciduous Forest  
Sunlit leaves, 1997

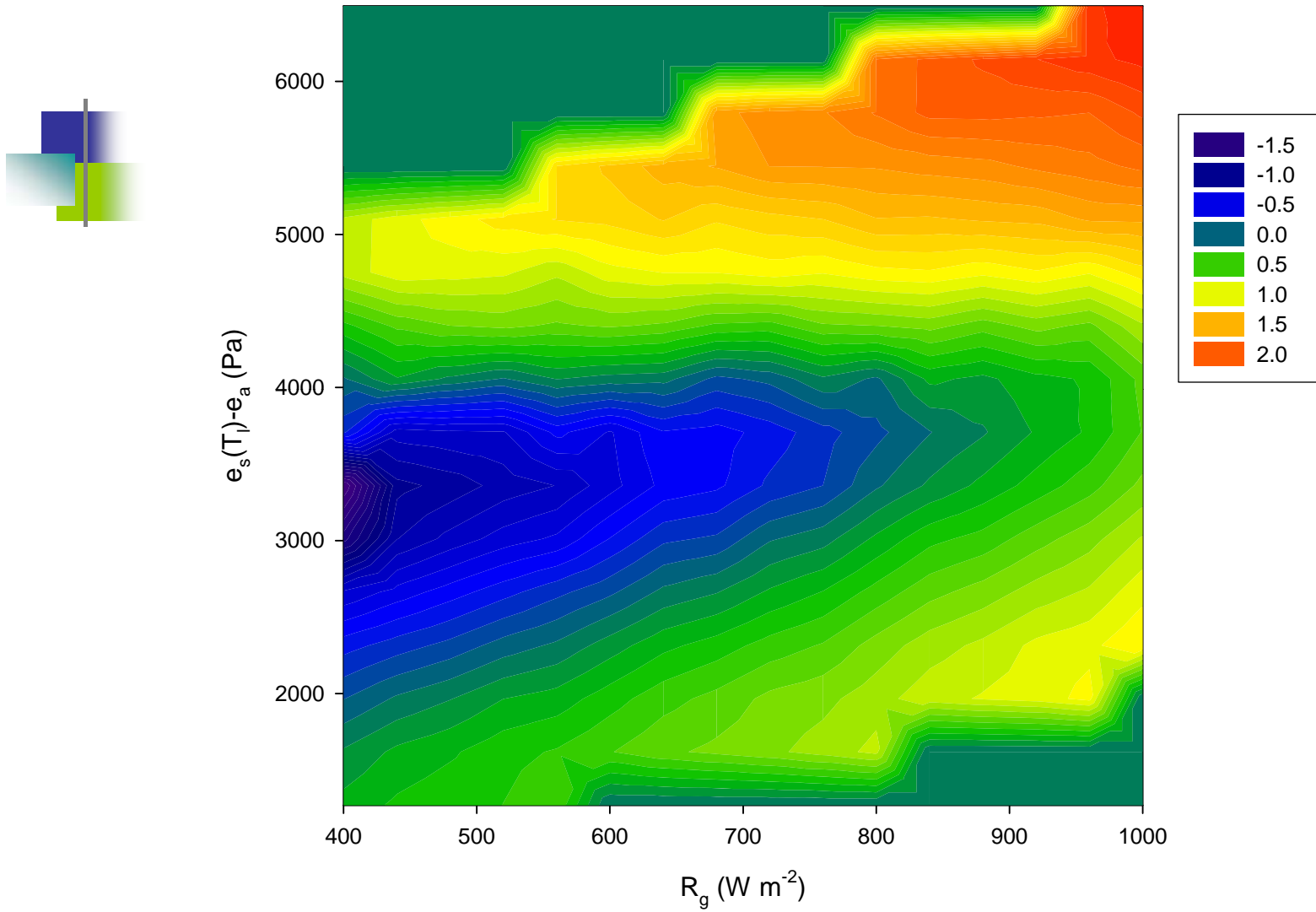


# Leaf Size, Temperature and Evaporation





$T_i(100\text{ mm}) - T_i(40\text{ mm})$



# Summary

- LE imposed is defined as the limit when boundary layer conductance goes to infinity. LE is independent of the net radiation balance!
- LE goes to zero when the boundary layer conductance approaches zero, for there is no physical mechanism to transfer water from the leaf to the atmosphere
- equilibrium evaporation is independent of stomatal conductance.
  - It results from the consequence of feedbacks between surface temperature, the temperature of the air and its water holding capacity.
  - It is also identical to the derivation based on the condition when  $\lambda E(R_n)$  is evaluated as  $g_a$  goes to zero.